Comparison of the Sentinel-1B Synthetic Aperture Radar with Airborne Microwave Sensors in an Extra-Tropical Cyclone

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Abstract—In winter 2017, the University of Massachusetts Amherst's Imaging Wind and Rain Airborne Profiler (IWRAP) was flown on a National Oceanic and Atmospheric Administration (NOAA) WP-3D Hurricane Hunter aircraft under the direction of scientists from STAR at NOAA/NESDIS over the North Atlantic ocean out of Shannon, Ireland. IWRAP is a dual-frequency, conically-scanning, profiling Doppler radar initially developed by the Microwave Remote Sensing Laboratory at the University of Massachusetts Amherst that is routinely installed on the NOAA WP-3D research aircraft. The flight on February 6, 2017 targeted a region of high winds (greater than $30\,\mathrm{m\,s^{-1}}$) that was also observed by the Sentinel-1B satellite's synthetic aperture radar. Sentinel-1B was configured to observe in extended wide swath mode in both VV- and VH-polarizations while the IWRAP C-band radar was configured to measure all of VV-, VH-, and HH-polarizations. IWRAP and Sentinel-1B VV and VH NRCS at the same Earth-incidence angle along the flight path match reasonably well during the entire flight, but some additional trends between aircraft and satellite can be observed. IWRAP VV-polarized NRCS generally match the CMOD5.h geophysical model function (GMF), suggesting errors in the Sentinel-1B processing chain.

Index Terms—C-band, ocean winds, cross-polarization, normalized radar cross-section, scatterometry.

I. Introduction and Experiment Description

SYNTHETIC APERTURE RADAR (SAR) observations of sea-surface backscatter are becoming popular for retrieving wind speeds from tropical cyclones because of their high spatial resolution and cross-polarization capabilities [1], despite the inherent challenges in retrieving wind direction. It is also difficult to calibrate and validate these instruments for use at all antenna elevation angles and high wind speeds because of

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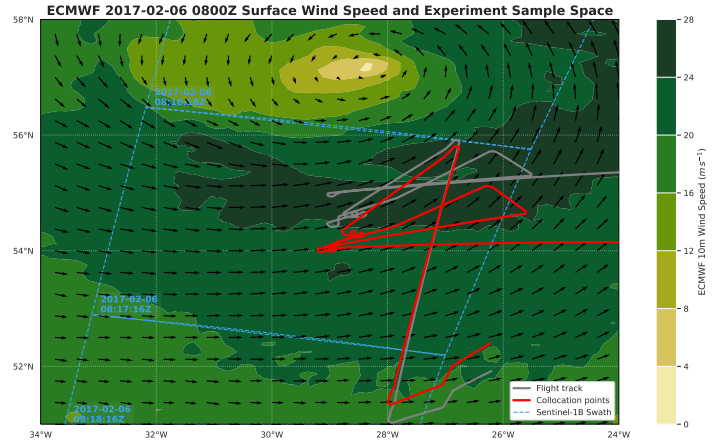Digital Object Identifier TBD

Figure 1. Flight track of the NOAA WP-3D (gray) and the collocation points with Sentinel-1B (red) on top of ECMWF wind contours and vectors. Sentinel-1B acquisition areas are shown as large trapezoids with the acquisition time, approximately 0817Z for most of the collocated points, in the bottom left corner. The Sentinel-1B nadir is to the right of the acquisition swath in the image, so the Earth-incidence angles increase from right to left.

their sparse data acquisition, but some satellite intercomparison is possible [2]. In particular, the very low backscattered cross-polarized signal imposes a severe constraint on the satellite noise estimate and correction. The Ocean Surface Winds Team from the Center for Satellite Applications and Research (STAR) at the National Oceanic and Atmospheric Administration (NOAA) National Environmental Satellite, Data, and Information Service (NESDIS) routinely performs airborne flight experiment campaigns for the purpose of collecting remotely-sensed and *in-situ* data in order to calibrate, validate, and improve the understanding of measurements from spaceborne sensors. The primary purpose of this manuscript is to evaluate the performance of the cross-polarized measurements from the Sentinel-1B SAR across subswaths in a rain-free high-wind environment. This has implications for satellite intercomparison with Sentinel-1B, wind speed retrievals from Sentinel-1B, and development of future cross-polarized geophysical model functions. We also identify areas of improvement for the calibration of Sentinel-1B.

On February 6, 2017, the authors directed a flight of a NOAA WP-3D over an extratropical cyclone in the North Atlantic while the Sentinel-1B SAR sampled the same region. Wind speeds were in excess of $30\,\mathrm{m\,s^{-1}}$, providing an opportunity to compare normalized radar cross-section (NRCS) from spaceborne SAR with aircraft measurements over a large field of near-hurricane-

force winds. An additional benefit to this storm in particular, and extratropical cyclones in general, is that the environment was cold and dry. There was no liquid precipitation under the aircraft, according to onboard instrumentation. These conditions greatly reduce the likelihood of enhanced surface roughness due to rain and create a stable environment for comparing the SAR with airborne sensors over a few hours. Fig. 1 shows the flight track over the European Centre for Medium-Range Weather Forecasts (ECMWF) wind vectors from 0800Z. Sentinel-1B acquisition areas are shown as large trapezoids.

The Sentinel-1 mission is a constellation of two polar-orbiting satellites, each carrying a C-band (5.405 GHz) synthetic aperture radar, and is operated by the European Space Agency for the operational needs of the Copernicus program. The Sentinel-1 C-band SAR supports operation in single polarization (HH or VV) and dual polarization (HH+HV or VV+VH) and have four exclusive acquisition modes. During the NOAA flight experiment on 6 Feb 2017 it was operated in extended wide swath (EW) mode, which is 400 km wide and observes the Earth at Earth-incidence angles (EIAs) from 17° to 45°. EW mode captures five sub-swaths using Terrain Observation by Progressive Scans SAR (TOPSAR). With this technique, in addition to steering the beam in range as in ScanSAR, the beam is also electronically steered from backward to forward in the azimuth direction for each burst [3]. The data presented here were from the two subswaths that correspond to the lowest EIA (i.e., near-range—approximately 16° to 35° incidence, depending on orbit altitude). We performed noise correction and averaging on the Sentinel-1B data. The final pixel spacing is 1 km but the resolution is 3 km.

The NOAA WP-3D carried the Imaging Wind and Rain Airborne Profiler (IWRAP) [4], [5], a C- and Ku-band scatterometer system designed and developed by the University of Massachusetts Amherst, and the Stepped Frequency Microwave Radiometer (SFMR), a C-band microwave radiometer. Though SFMR was developed for use in tropical environments, cold-weather corrections have been applied so that they match GPS dropwindsondes better [6].

IWRAP is a Doppler radar that was originally designed to support a conical scan at multiple incidence angles and polarization configurations. An internal calibration loop allows for tracking of receiver gain and transmit power during the flight. Since winter 2015, the C-band radar has been configured for a fan-beam antenna on loan from ESA [7], [8]. This antenna has a better cross-polarization isolation (greater than 40 dB isolation across the main lobe) than the C-band IWRAP pencil-beam phased array, allowing for measurements of cross-polarized NRCS at low wind speeds. It is statically mounted and, since summer of 2015, pointing 25° off nadir and towards the right of the aircraft. The IWRAP C-band radar transmits a pulse-compressed waveform (i.e., chirp) with a 4 MHz bandwidth at the antenna's design frequency of 5.3 GHz. The ESA antenna on IWRAP has an azimuthal (along the direction of the aircraft fuselage) beamwidth of approximately 5° across the main lobe. This equates to approximately 80 to 125 m on the surface for 20° to 30° EIA—the range used in this manuscript—but the spatial resolution was coarsened to 1 km in the direction of the aircraft flight track to reduce noise levels and better match the spatial resolution of the Sentinel-1B SAR. IWRAP has a range gate spacing of

30 m, but the cross-track resolution (along the direction of the aircraft wings) on the ground is approximately 90 to 60 m across the same EIA range. So even though IWRAP has a comparable resolution along the flight track, it has a much smaller cross-track resolution. Both V and H polarizations are available on transmit and receive and, for this experiment, IWRAP measured NRCS at VV, HH, and VH. With the exception of the C-band antenna mount angle, the configuration is the same as in [5]. While the Ku-band radar was operational at two incidence angles for VV and HH, the C-band radar is the radar of interest in this manuscript. The C-band fan-beam antenna allows selection of NRCS from approximately 15° to 45° EIA at level flight and the typical altitude of 2 to 3 km.

The SFMR has a 3 dB beamwidth of 20° to 28° and it takes approximately 4 s to step through all frequencies [9]. For the flight on 2017 Feb 6, this beamwidth equated to a cross-track swath of 750 to 1050 m, which is close to the spatial scale of the Sentinel-1B measurements used and the alongtrack averaging scale of IWRAP. During the integration time, the aircraft traveled 0.5 km on average. Using the same beamwidth values as the cross-track estimates, this yields an alongtrack averaging distance of 1250 to 1550 m. These small differences and variations in spatial scale are not expected to change the analysis since surface wind speeds are only used to adjust Sentinel-1B VV NRCS and to check for biases in IWRAP. As will be described in the next section, adjustments to correct for different azimuth look angles are made only to the Sentinel-1B VV NRCS. All VH conclusions are independent of wind speed.

All sensors have similar spatial resolutions, with the exception of the IWRAP cross-track resolution. This will result in noisier IWRAP data, but otherwise do not affect the analysis.

The remainder of this paper is structured to show an elevation angle-dependent NRCS difference between Sentinel-1B and IWRAP, especially as a function of subswath. Elevation angle was chosen for most of the discussion because the data suggest a radar calibration problem rather than a geophysical phenomena. This is more naturally expressed for a satellite as a function of elevation angle.

Section II describes how the sensors were aligned in time and space to measure the same ocean-surface phenomena; Section III details the IWRAP calibration procedures; and Section IV presents and interprets intercomparison results.

II. Collocation Methodology

Collocation of airborne sensors with spaceborne sensors in extreme environments is challenging due to the speed difference of the two platforms and the movement of the storm. While it took minutes for Sentinel-1B to image the area, the WP-3D took hours, during which the location, direction, and magnitudes of the winds initially sampled by the SAR changed. In order to account for the storm displacement, we used storm-relative positions to collocate the two instruments.

Aircraft locations and times were first grouped into 1 km alongtrack cells. Then we estimated the storm center based on the minimum wind speed location from spatial and temporal interpolations of Global Data Assimilation System (GDAS) model runs between 0700Z and 1200Z (the approximate time
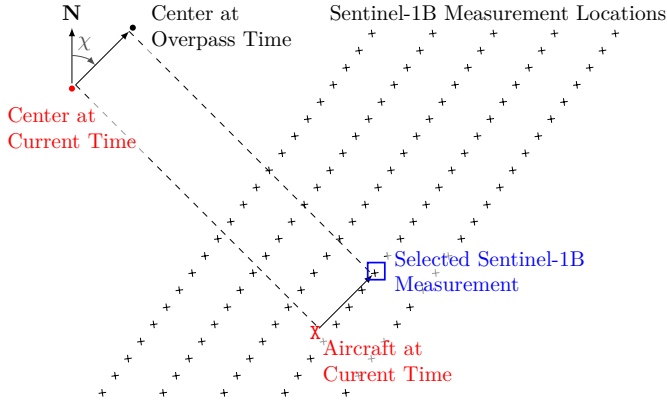
Figure 2. A schematic diagram showing how the Sentinel-1B measurement location is selected. The aircraft location is translated in the same bearing χ and distance (in km) that the storm center is from the current time to the satellite overpass time. The nearest Sentinel-1B measurement to the resulting location is selected for collocation. All IWRAP alongtrack cells are passed through this procedure.

the aircraft was collecting data). The models provide adequate estimates for the storm center but are not used as wind direction ground truth for either sensor, as we describe below. The track of the aircraft was offset in the direction of storm motion before the overpass and in the opposite direction after the overpass, by the distance difference between the estimated storm centers at aircraft-measurement time and satellite-measurement time. The nearest Sentinel-1B measurement to the resulting location is selected for collocation. Fig. 2 shows a schematic diagram of this location translation procedure.

Since the IWRAP C-band radar has a range of incidence angles to choose from, along each point in the collocated flight track the IWRAP measurements with an EIA within $\pm 1°$ of that of the selected Sentinel-1B observation was selected. The NRCS measurements were then averaged into the same alongtrack cells that the aircraft positions are. In the following analysis, we assume both the magnitude and direction of the winds do not change significantly. Though there were most certainly some small-scale changes between the times the aircraft and satellite instruments sampled the storm, they do not affect the conclusions drawn from the comparison. While this is not possible to prove unequivocally, the auxiliary data shown below support this assumption. The dropsonde wind speeds and directions match the corresponding GDAS-interpolated 10 m winds relatively well. The GDAS values are interpolated from the model runs at 0600Z and 1200Z, which is a greater time span than the aircraft flight. This suggests stability in the surface wind field, the primary driver of the NRCS shown in the rest of the paper. Finally, NWS/NCEP OPC (Ocean Prediction Center) Atlantic surface analyses valid at 0600Z and 1200Z[1] only show a difference in central pressure of the storm of 1 mbar, which is a negligible change.

[1]See https://www.ncei.noaa.gov/thredds/catalog/ncep_charts_new_charts_ocean_anl/201702/20170206/catalog.html

III. Calibration

For most of the aircraft flight time, Sentinel-1B and IWRAP were not observing the ocean at the same azimuth in the same direction, which results in differences in the measured NRCS that are not due to instrument or processing differences. We compensated for this by using a geophysical model function (GMF) to predict differences between the two wind-relative observation angles, and have applied the difference to the Sentinel-1B NRCS. Note that this does not remove the wind direction effect on the observed NRCS. When showing the Sentinel-1B NRCS after compensating for azimuth look angle in the following results, they are labeled "ALAC." A more detailed description of this processing step follows.

After averaging IWRAP NRCS in the alongtrack dimension, wind directions retrieved from the Sentinel-1B images [1] were also collocated with the alongtrack cells. Both the IWRAP C-band azimuth and the Sentinel-1B azimuth were subtracted from these wind directions to obtain a wind-relative azimuth. Using these relative wind directions, the mean incidence angles of each sensor in the alongtrack bin, and SFMR wind speeds as the input to the CMOD5.h GMF [10], modeled NRCS ratios of the two sensor geometries were computed and applied to the Sentinel-1B VV NRCS. In other words, for each alongtrack bin the quantity

$$\Delta\sigma^0_{\text{ALAC}} = \frac{\sigma^0_{\text{CMOD5.h}}(U_{10N}, \phi_{\text{IWRAP}}, \theta)}{\sigma^0_{\text{CMOD5.h}}(U_{10N}, \phi_{\text{S1B}}, \theta)} \qquad (1)$$

was computed for NRCS in linear units and multiplied with the Sentinel-1B NRCS, where U_{10N} is the 10 m equivalent-neutral wind speed reported by SFMR, ϕ is the wind-relative azimuth angle of each sensor, and θ is the incidence angle observed by both sensors. This has the effect of rotating the Sentinel-1B observation angle to match IWRAP and leaves any non-geophysical effects intact. Since there is not much wind direction dependence at cross-polarization [5], this step was not performed for the VH NRCS.

Ground calibration of IWRAP was performed before and after the flight experiment campaign in order to measure the receiver gain and transmit power levels. Because this does not include the antenna, a final, constant calibration offset is determined from the sea-surface NRCS. The calibration offsets change between installations due to radar reconfiguration, maintenance, and the simple fact that disconnecting cables for platform deintegration or ground calibration creates a slightly different measurement path.

Wind speed retrievals from SFMR were performed using cold-weather corrections to the tropical-weather GMF [11] to obtain reference wind speeds collocated with IWRAP. SFMR wind speeds from 22 to 24 m s^{-1} were selected along with IWRAP NRCS measurements at incidence angles of 29° to 31° from this flight. The CMOD5.h GMF was used with wind directions as above to compute the true NRCS. The incidence angle range is within the valid range of CMOD5.h (which inherits its limits from CMOD5.n [12]) and the C-band fan-beam of IWRAP during level flight. The wind speeds were chosen such that they would target a large area at the beginning of the collocation region and be in a valid range for both SFMR and CMOD5.h. The median of the differences between the GMF and measured NRCS at VV-polarization was added to the IWRAP VV NRCS.

The HH calibration adjustment is estimated by computing the VV-only CMOD5.h and applying the wind-speed- and EIA-dependent polarization ratio from [13].

Assuming these small, constant calibration offsets are from the antenna connections, each adjustment is a result of the electromagnetic waves propagating both into and out of each polarization port. Therefore, the cross-polarization calibration adjustment is the sum (in dB) of half of the co-polarized offsets (the geometric mean in linear units). We do not rely on a VH GMF for calibration because none are mature enough in our opinion, and it is more appropriate to correct for the source of the error (in this case, the V- and H-polarized antenna connections) than the resulting behavior (VH NRCS).

All of these calibration adjustments are on the same order of magnitude as in previous years (e.g., [5], [14]). HH-polarization results are not shown here since Sentinel-1B was not measuring HH NRCS at the time.

This procedure neither causes the IWRAP NRCS to track SFMR wind speeds nor does it affect the incidence angle dependence; it only ties the means of the IWRAP measurements in the small wind speed range to a fixed value. Thus the independence of the two systems are retained. However, there still may remain a constant offset to the true NRCS in both VV and VH.

Polarization mixing is the effect that occurs when rotating the antenna's electric field away from perpendicular or parallel to the plane of the ocean. Some amount of HH-polarized NRCS is "mixed" with the VV measurements, and co-polarized NRCS likewise contaminate cross-polarized measurements. The correction for this phenomena is most noticeable in the VH measurements. In the following results, we have calculated and removed the contribution to measured VH NRCS due to polarization mixing using the aircraft's pitch, the antenna mount angle, and (3) in [5]:

$$\sigma_{VH}^{0}{}' = \sigma_{VH}^{0} + \left(\sigma_{VV}^{0} + \sigma_{HH}^{0}\right)\sin^2\gamma\cos^2\gamma, \qquad (2)$$

where the prime symbol indicates the measured quantity and γ is the negative value of the aircraft pitch angle. The pure NRCS is simulated using the same co-polarized geophysical model functions used for calibration. This adjustment is less than 0.1 dB for more than 95 % of the VH data.

IV. Results and Discussion

Fig. 3 shows the IWRAP and Sentinel-1B NRCS as a function of time for VV-polarization (top panel) and VH-polarization (bottom). In the bottom part of each panel, the SFMR 10 m equivalent-neutral wind speed is shown along with the same wind speed estimates from GPS dropwindsondes [15] and GDAS. As the dropsondes measure a profile, an averaging and scaling method (WL150) has been applied to estimate the equivalent-neutral wind speed at 10 m. This algorithm is described in [5]. Some data appear missing due to being outside the swath (at approximately 1000Z; see also Fig. 1) or during large-angle turns or 360° orbits (at approximately 0905–0910Z and 1045Z), where the SFMR retrievals are not fully validated.

The dropsonde surface wind speed estimates are slightly lower than SFMR wind speeds in general. A few report much lower speeds than SFMR, but this has been observed in other
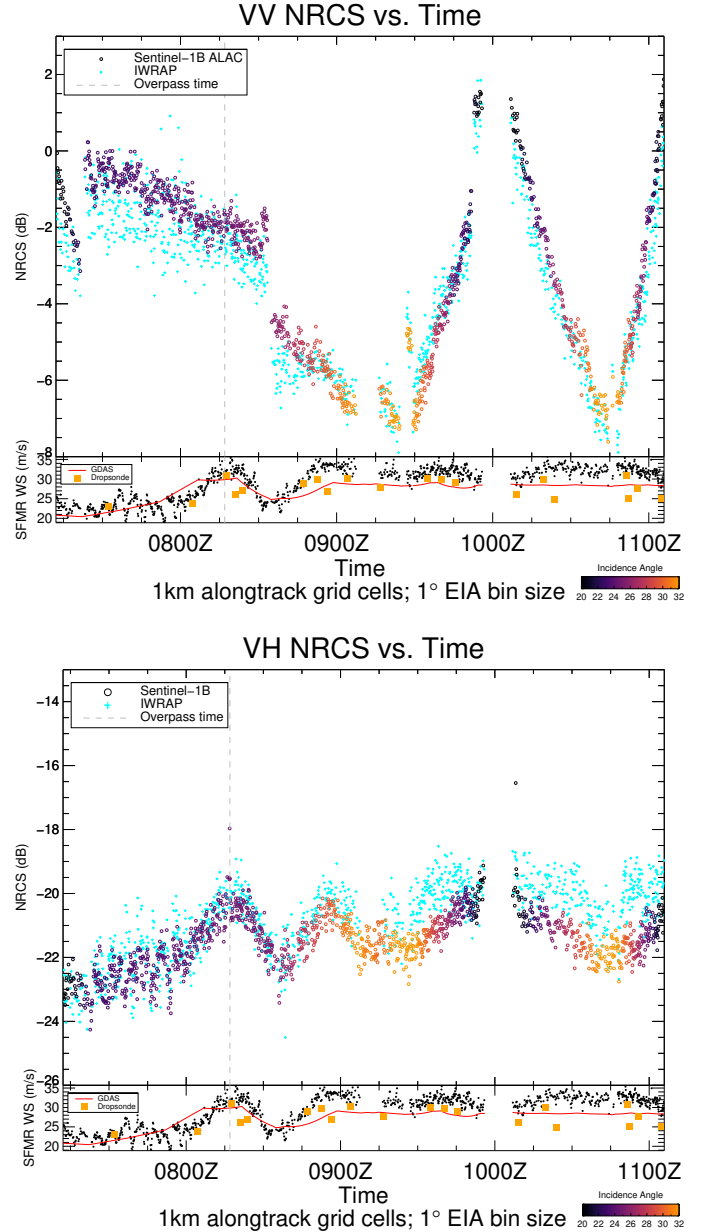

Figure 3. NRCS (in dB) as a function of time during the flight on 6 Feb 2017. The upper panel shows VV-polarization and the lower panel shows VH-polarization. The main part of each panel shows the IWRAP NRCS as crosses and the Sentinel-1B NRCS as empty circles colored by the incidence angle. Within each panel, the lower part shows the wind speed retrieved from SFMR as well as the surface wind speed estimation from dropsondes.

comparisons using a relatively small number of dropsondes [5] and do not necessarily indicate errors in SFMR retrievals. It is possible that SFMR wind speeds are too high due to imperfect cold-weather corrections. However, the conservative approach for the present analysis is to use the SFMR wind speeds as they are. Reducing the wind speed used in (1) increases the adjustment due to ALAC. This increases the slope of the VV sensor ratios (top panels of Figs. 4 and 5) and results in a greater difference between subswaths.

The dropsondes appear to align with the interpolated GDAS wind speeds relatively well, but GDAS does not adequately represent the higher gradients. This is not only apparent com-

pared to SFMR wind speeds, but also with the two VH NRCS measurements.

As can be seen in Fig. 1 and the coloring of Fig. 3, the track of the aircraft proceeded generally towards an increasing incidence angle until approximately 0930Z. At this time, the aircraft turned East, exits the swath near 1000Z, and turned back West again. Outside of the swath is below the lowest EIA observed by Sentinel-1B, so during this Eastbound leg the VV NRCS can be seen to have increased with time as EIA decreases. After having reached the Westernmost point at approximately 1045Z, the aircraft turned East again. From 0900 to 1100Z the wind speed remained relatively constant, so changes in the VH NRCS measured by both sensors during this time span are due to the weak EIA sensitivity.

The VV-polarized NRCS are not obviously correlated with wind speed, shown in the lower part of each panel of Fig. 3, but are strongly influenced by EIA and wind-relative look angle. For example, from approximately 0715Z to 0830Z the EIA changed from $23°$ to $25°$ and the wind-relative azimuth changed from $180°$ to $140°$. There is sensitivity of VV NRCS to wind speed at these speeds, but it is masked by the stronger sensitivity to EIA.

Though CMOD5.h may not be absolutely precise at these low incidence angles (due to limitations in the development of CMOD5.n; see below for further discussion), its behavior is useful for understanding the VV response to changes in EIA and azimuth in general. For the values of wind speed, EIA, and wind-relative azimuth of the sensors during this time, CMOD5.h predicts 1.5 to 2 dB of NRCS decrease. This is nearly the same amount of change observed by both IWRAP and Sentinel-1B. Since differences with IWRAP due to azimuth angle have been removed from the Sentinel-1B co-polarized data, any remaining difference between IWRAP and Sentinel-1B may be primarily attributed to calibration errors in IWRAP or Sentinel-1B.

In contrast to VV, the VH NRCS measurements from both sensors follow the same trend as the SFMR wind speed and are weakly influenced by EIA. The VH NRCS tracks the changes in wind speed despite the changes in EIA and wind-relative azimuth.

Note that two elevated Sentinel-1B VH samples at the collocation time and near 1010Z, visible near the -18 dB level in the bottom panel of Fig. 3, do not seem to be from the ocean surface backscatter but instead from the aircraft. Since these don't appear in the co-polarized channel, it isn't obvious whether they are from interaction with the body of the airplane or some other active transmitter on the airplane. The aircraft interference with the Sentinel-1B measurement exists across the satellite track of subswath 1, so it is the closest collocated measurement in two different locations.

Fig. 4 shows the ratio of IWRAP to Sentinel-1B NRCS in linear units as a function of Sentinel-1B antenna elevation angle. The individual comparison points are colored by SFMR wind speed to show that any trend is not due to wind speed contributions. Fig. 5 is the same but for the data prior to ESA-reprocessing made available only to the Expert Support Laboratory working on Sentinel-1 in Spring 2018—in other words, the data available to the public. A linear fit was performed to the means.

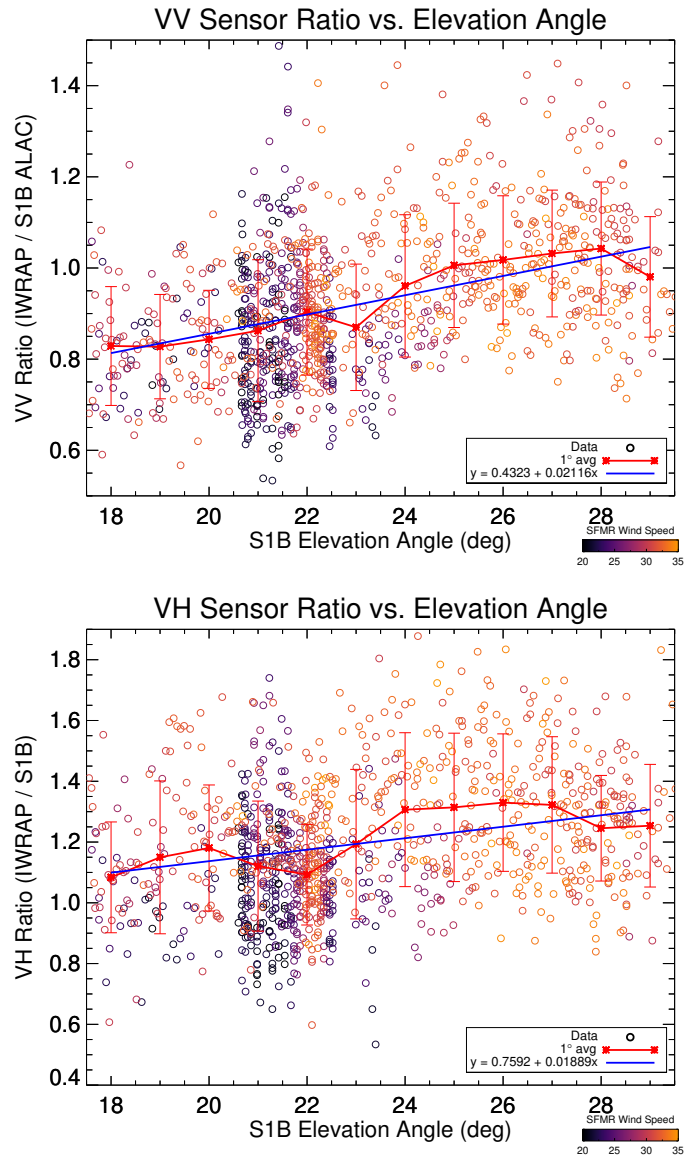Vertical offsets from a ratio of 1 indicate that IWRAP NRCS is

Figure 4. NRCS ratio (in linear units) between IWRAP and Sentinel-1B as a function of the Sentinel-1B elevation angle during the flight on 6 Feb 2017. The upper panel shows VV-polarization and the lower panel shows VH-polarization. The empty circles show the raw data colored by SFMR wind speeds and the asterisks connected by a solid line are means and sample standard deviations within $1°$ elevation angle bins. The solid line shows a linear fit to the means, with the equation given in the legend.

larger (positive offset) or smaller (negative offset) than Sentinel-1B NRCS. There is a slight trend in both VV and VH ratio, and the two slopes are nearly the same. This was not just an effect of time difference; as can be seen in Fig. 3, the elevation angle being used varied throughout the flight independent of time. There is not an obvious surface wind speed dependence, as retrieved from SFMR.

To rule out a bias in the IWRAP NRCS with incidence angle (due to knowledge of the gain pattern, surface area calculation, or errors resulting from incorrect antenna attitude), we show the ratio of modeled NRCS to both IWRAP and S1B NRCS as a function of Sentinel-1B elevation angle in Fig. 6. The NRCS GMF chosen was CMOD5.h; we used SFMR retrievals as the wind speed input and the same wind direction described
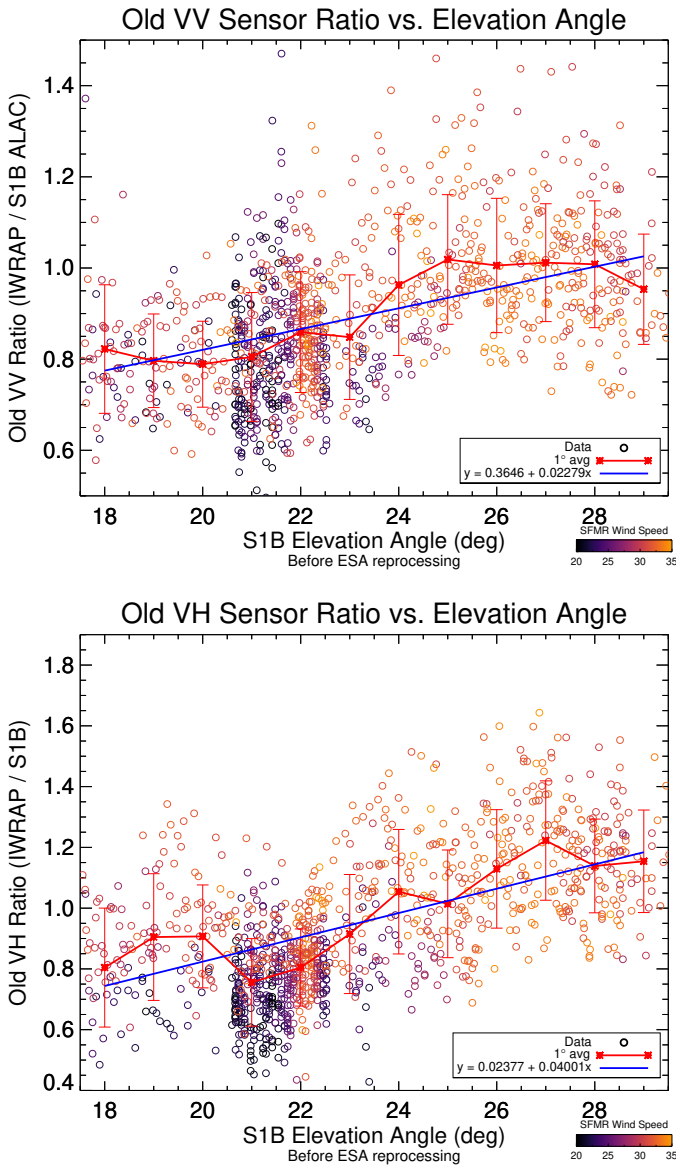
Figure 5. NRCS ratio (in linear units) between IWRAP and Sentinel-1B as a function of the Sentinel-1B elevation angle during the flight on 6 Feb 2017 prior to the ESA reprocessing in Spring 2018. The upper panel shows VV-polarization and the lower panel shows VH-polarization. The empty circles show the raw data colored by SFMR wind speeds and the asterisks connected by a solid line are means and sample standard deviations within 1° elevation angle bins. The solid line shows a linear fit to the means, with the equation given in the legend.
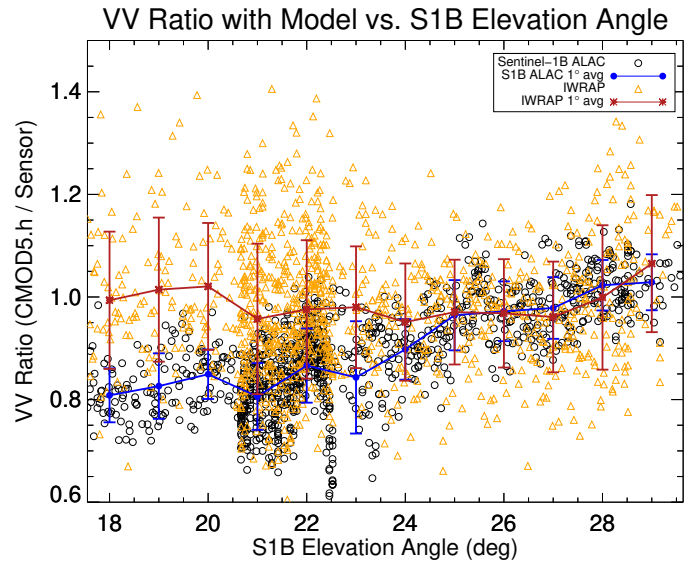

Figure 6. VV-polarized NRCS ratio (in linear units) between the CMOD5.h GMF and each sensor as a function of the Sentinel-1B elevation angle during the flight on 6 Feb 2017. Empty circles show the Sentinel-1B data and empty triangles show the IWRAP data. Filled circles show the 1° average of the Sentinel-1B data and asterisks show the same for IWRAP. Surface wind speeds were from SFMR and wind directions were retrieved from the Sentinel-1 image.

GMF work [16] have observed and compensated for uncertainty in CMOD5.n at low incidence angles (for example, see Fig. 3 of [16]). Considering this, the NRCS measured by IWRAP is within an expected margin of error of CMOD5.h. The IWRAP data at low incidence angles (18 to 20°) suggest that the low-incidence angle component of CMOD5.n should be revisited, as has been done with CMOD7 [16].

We do not use CMOD7 here because it was not developed specifically for high winds as CMOD5.h was and it uses stress-equivalent wind speeds as the surface wind speed. This is not the same wind speed as the SFMR or dropsonde reports. Since it only causes a small vertical shift in the NRCS of Fig. 6 (not shown), we use CMOD5.h for this figure and the ALAC.

There is definitely a strong trend in the Sentinel-1B VV NRCS ratio, while the IWRAP VV measurements show little difference compared to the model. Fig. 6 shows that the VV observations in Figs. 4 and 5 are not due to a bias in IWRAP. A larger scatter in the IWRAP data is noticeable despite the alongtrack averaging on the same order as the Sentinel-1B averaging. This is due to the smaller cross-track size.

Fig. 7 shows the VH NRCS as a function of incidence angle for IWRAP and Sentinel-1B. To make a comparison of the incidence angle dependence of the VH NRCS between the two sensors, the NRCS data where SFMR wind speeds are higher than $30 \, \mathrm{m \, s^{-1}}$ have been averaged in 1° EIA bins and plotted as filled circles. The darker colored points were measured at lower surface wind speeds. The strong wind speed sensitivity can be observed in the difference between these and the values above $30 \, \mathrm{m \, s^{-1}}$.

Though the connected points of Figs. 6 and 7 suggest a smooth transition between Sentinel-1B subswaths, this may not be the case. The subswath transition occurs at approximately 24° elevation. A closer look at Fig. 6 reveals a good match (ratio of 0.9 to 1) between modeled and measured NRCS above 24° and a

in Section III. Comparisons with IWRAP alongtrack-averaged data and 1° means are shown as empty triangles and asterisks, respectively. The same is shown for Sentinel-1B as empty circles and filled circles. A value below 1 indicates the sensor is reporting higher NRCS than modeled.

One may expect that IWRAP NRCS be nearly 1 in the mean in Fig. 6 due to the calibration procedure. While it is close (within 1 sample standard deviation), it appears to be a little low. Recall that IWRAP was calibrated at 30° incidence (±1°), which corresponds approximately to the 25.5 to 27.5° elevation angle range; the CMOD5.n (and, by extension, CMOD5.h) GMF at incidence angles below 27° was developed using a different sensor than for the higher incidence angles. Recent
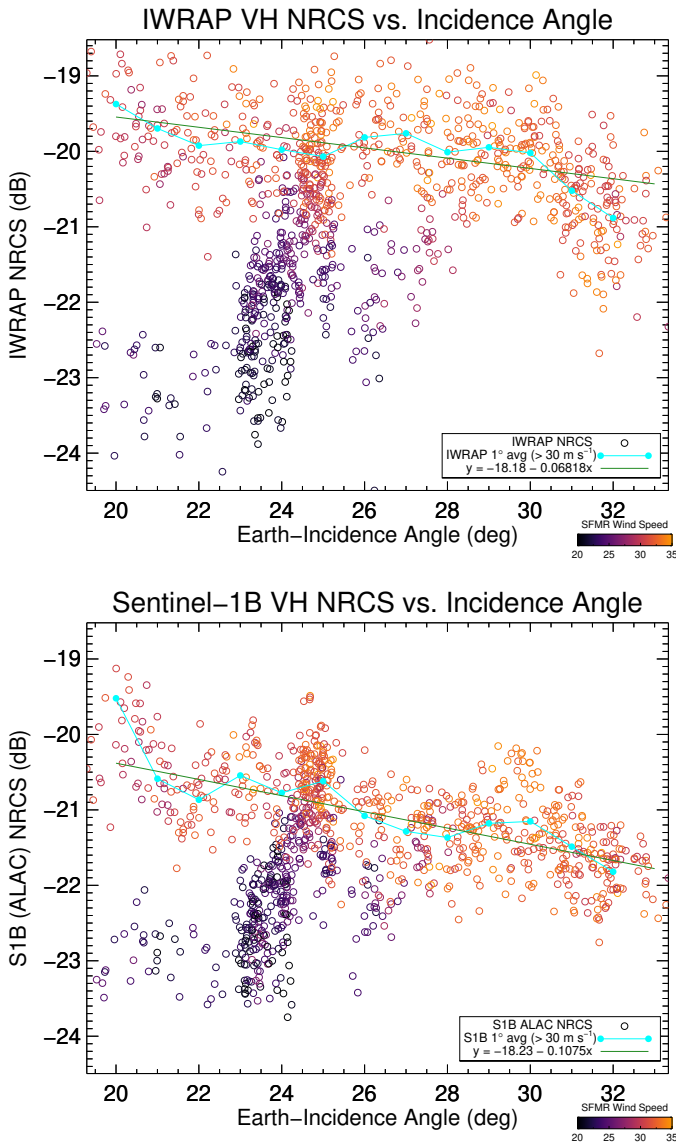
Figure 7. VH NRCS vs. Earth-incidence angle for the two sensors from the flight on 6 Feb 2017. The upper panel shows IWRAP and the lower panel shows Sentinel-1B. The points are colored according to SFMR wind speed. Filled circles connected by a solid line are the means of the data points where the SFMR wind speed is higher than $30 \, \mathrm{m \, s^{-1}}$ in 1° elevation angle bins. The solid line shows a linear fit to the means, with the equation (fit in dB) given in the legend.

clear lower ratio (below 0.8) below 23°. It is not clear whether or not the NRCS in subswath 2 are mis-aligned, but it is apparent that the measurements in subswath 1 (low elevation angles) are biased high.

As noted in Section III, the absolute calibration level of the IWRAP VH channel may not be correct. Based on Fig. 6, we expect the higher incidence angles to match better in power. Comparing with Fig. 4, the IWRAP VH channel may be 1.3 dB too high. However, the relative calibration is reliable and the trends shown in Fig. 7 can be analyzed in light of these facts. Despite the same geophysical scene and the same decreasing trend of VH NRCS with increasing EIA, Sentinel-1B clearly shows a stronger dependence of VH NRCS on EIA. This is consistent with our conclusions about VV NRCS drawn from Fig. 6: that the Sentinel-1B subswath 1 measurements are too

high compared to IWRAP measurements.

Some papers using RADARSAT-2 SAR data have concluded that no incidence angle sensitivity exists for VH NRCS [13], [17]. However, both sensors show a weak VH NRCS dependence on EIA, which is consistent with other studies (e.g., [1], [18]–[20]) and theory (e.g., [21], [22]). It is also consistent with prior IWRAP-based findings [23], though to a lesser degree than [5]. The IWRAP VH data in Fig. 7 show a more subtle slope than those from Sentinel-1B.

V. Summary

The data collected from the NOAA WP-3D flight experiment on 6 Feb 2017 contributes to at least two of the objectives of the NOAA Ocean Winds flight experiment program: development and improvement of space-based ocean surface vector wind products, and testing of new remote sensing technologies for future satellite missions.

The data presented here further demonstrate the utility of having both VV- and VH-polarized channels on a scatterometer or SAR. With sensitivities to different ocean parameters, both polarizations further constrain wind vector retrievals. The second-generation MetOp scatterometer (MetOp-SG SCA) will have both VV and VH capabilities, and this collocated dataset clearly demonstrates the strong sensitivity of the VH NRCS signal to wind speed from both aircraft and satellite sensors. The data also show the relative insensitivity of VH to wind direction, a feature that should prove useful for wind speed retrievals from the scatterometer on MetOp-SG. Since only the antenna perpendicular to the flight path on each side of the spacecraft is planned to have the cross-polarized channel, this insensitivity may reduce the need for multiple looks at a wind vector cell from different directions. This analysis also confirms that there is no or very little sensitivity to the incidence angle for VH NRCS.

In the course of writing this manuscript ESA reprocessed the Sentinel-1B data presented here for use by the Expert Support Laboratory working on Sentinel-1. A significant trend of S1B VH NRCS with elevation angle, compared to IWRAP NRCS, was improved and a step at subswath edges was reduced. However, VV NRCS largely remained the same. The VH NRCS still show a small bias in one of the subswaths. The choice of ocean calibration model functions impact the absolute level of the IWRAP VH NRCS. There may be a trend in the whole range of antenna elevation angles or it may be limited to a discontinuity at the subswath boundary.

The VV NRCS show a similar bias or trend as the VH do, but inaccuracies in surface wind speed or direction would lead to a different ratio between Sentinel-1B and IWRAP. There is better agreement with IWRAP and modeled VV NRCS generated from SFMR wind speeds and Sentinel-1B wind directions in subswath 2, which suggests that there are processing-related problems in subswath 1.

Overall, this paper confirms that getting both calibrated VH NRCS from satellites and intercalibrated measurements from different remote sensing systems remains a challenge. A careful analysis and monitoring of the Sentinel-1 SAR across the swath over permanent scatterers, such as a rain forest (similar to [24]) is certainly needed to fully characterize the antenna gain pattern.

The same is true for the noise annotated in the Level-1 products and are confirmed by the recent updates in the ESA Sentinel-1 Level-1 processor.

Interestingly, this study reveals that the inconsistency between IWRAP and Sentinel-1 or between CMOD5.n (VV GMF) and Sentinel-1 are more pronounced at low incidence angles where no ASCAT data exist to define the VV GMF. This inconsistency will impact the wind speed retrieval derived from VV NRCS at low incidence angles from Sentinel-1. This point also suggests a closer look be taken at the VV GMF that is to be used for the next-generation MetOp-SG SCA instrument, which will have lower incidence angles than the existing ASCATs.

Finally, regarding VH NRCS: if used for wind speed retrieval, the choice of the VH GMF may depend on the sensor. In particular, the VH GMF proposed from RadarSat-2 or Sentinel-1 data may not be appropriate for MetOp-SG SCA.

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